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1 **Basin-scale land use impacts on world deltas: human vs natural forcings**

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19 **Abstract**

20 A new global database of 86 deltas and river basins was analyzed to investigate the relative
21 importance of deforestation and land use changes versus natural forcings in determining long-
22 term total delta size. Results show that mean river flow and shelf slope were the most important
23 variables, whereas population density and sediment load had a much lower importance.
24 Deforestation and other variables related to land-use generally had a very small effect, but were
25 more influential in a subset comprising Mediterranean and Black Sea deltas. As most deltas have
26 developed over thousands of years, the much shorter-lived anthropogenic signals from
27 deforestation and other landscape perturbations have had only secondary impact on the total area
28 of deltas. Also, delta progradation is strongly influenced on sand deposition, whereas
29 anthropogenic impacts on sediment load have more often impacted mostly the finer sediment
30 being deposited offshore (prodelta deposits) or in the deltaic plain. These data disproves the
31 hypothesis that delta size and growth is strongly influenced by human forcings, particularly for
32 larger deltas, since Holocene delta building is mainly determined by natural forces. However,
33 humans are influencing the geomorphology of deltas, particularly over the last century when the
34 Anthropocene nature of deltas has become manifest. A more precise terminology is proposed to
35 clarify concepts such as “human-made”, “human-engineered” or “human-influenced” deltas.

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38 **Introduction**

39 Deltas are by definition sedimentary features, built by deposition of material eroded from river
40 catchments in conjunction with *in situ* organic soil formation. There are many anthropogenic
41 influences on sediment yield, including urbanization, deforestation, agricultural practices,
42 mining, and the retention of sediment by reservoirs (Syvitski et al., 2005). Not surprisingly,
43 therefore a growing number of studies have highlighted the anthropogenic alteration of fluvial
44 sediment flux to the coast (Walling, 2006), and concluded that some deltas have prograded much
45 faster or even formed during the late Holocene as a consequence of elevated sediment loads due
46 to human activities (Syvitski and Kettner, 2011; Maselli and Trincardi, 2013) (see Table S1 for
47 an expanded list of references). Deltas in Southeast Asia, the Mediterranean and Black Sea are
48 considered the paradigm of human influence on delta geomorphology and growth, because of
49 intense human activity within their basins over several millennia (Hori et al., 2001; Giosan et al.,
50 2012). The Ebro, Rhone, Po and Danube deltas have been highlighted as “man-made” (Maselli
51 and Trincardi, 2013; Guillén and Palanques, 1997).

52 The total size of a delta plain (subaerial delta), as defined by the extent of the delta fringe, is
53 primarily determined by the amount of coarser sediment fraction (fine sand) delivered by the
54 river, which in turn is dependent on river discharge (Syvitski and Saito, 2007; Giosan et al.,
55 2014). The interplay of fluvial input with marine waves and currents controls the partitioning of
56 sediment between subaerial and subaqueous portions of deltas (Swenson et al., 2005).
57 Consequently, increased sediment delivery is believed to result in a faster delta progradation and
58 larger delta size (Meade, 1996) and basins with long-term human-impacts (enhanced sediment
59 flux) might be expected to have deltas proportionally larger than less impacted ones. However,
60 the effect of increased fluvial sediment delivery on delta progradation depends on several factors,

61 including granulometry. Thus, a smaller grain size increases the fraction of sediment delivered to
62 the shallow marine area (the prodelta environment) and the extent of subaqueous delta
63 progradation relative to subaerial delta development (Swenson et al., 2005). Moreover, the ratio
64 of sediment delivery to river discharge differs among basins and is difficult to estimate (Walling,
65 1999; Notebaert et al., 2010). For example, large Asian river basins are often cited as examples
66 of low sediment delivery ratios (Walling, 2006; Allison et al., 1998), but some of them have high
67 human-induced sediment load (Saito et al, 2001; Shi et al., 2010). Such examples suggest that
68 there is an apparent contradiction between estimates of sediment fluxes near the source and those
69 at the sink, the so-called sediment delivery problem (Walling, 1999). Ultimately this is a question
70 related to the cumulative influence of the 7,000 year history of the Holocene, compared to the
71 recent history of the massive human-induced changes during the Anthropocene (Waters et al.,
72 2016).

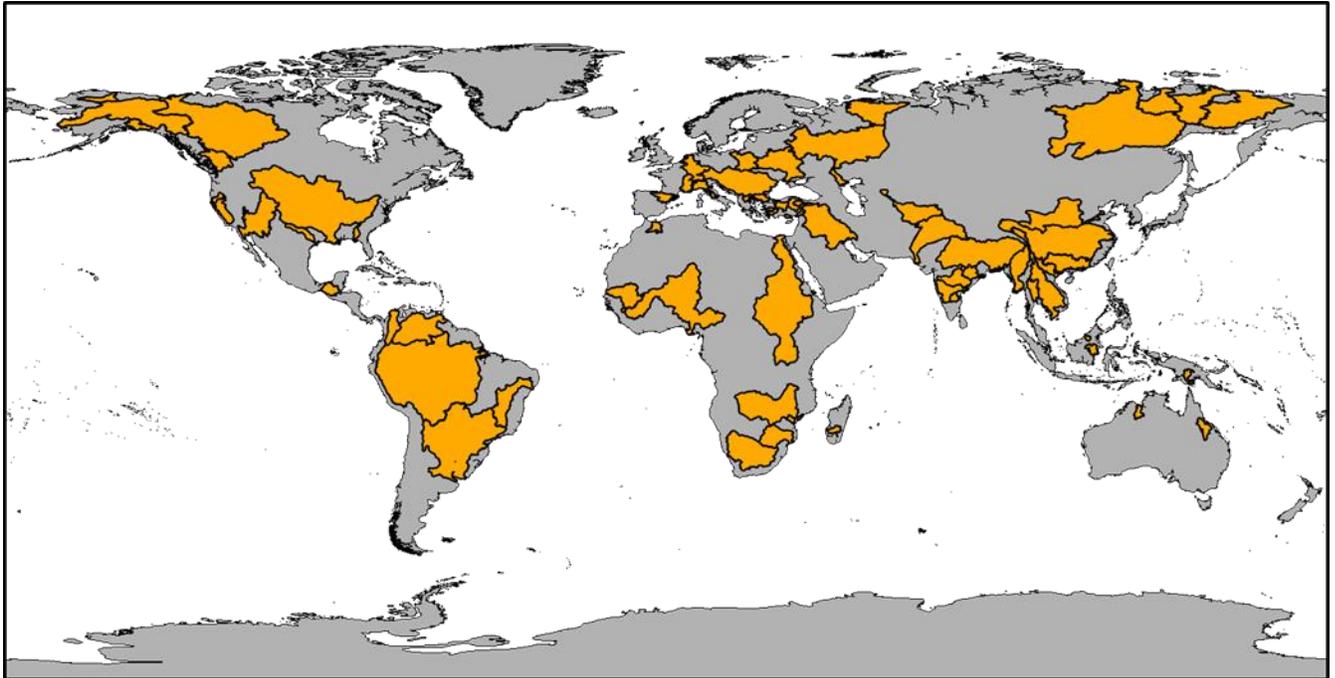
73 The mismatch between what is eroded from the basin, what is transported to the river mouth
74 and what is deposited in the delta can be attributed to several factors: a) deforestation and land
75 use change that do not always increase land erosion (Hamilton, 1987); b) most of the eroded
76 sediment is deposited close to its source (Wilkinson and McElroy, 2007); c) coarser sediment
77 transport involves a variable time-lag (10's to 1000's of years) from source to sink depending on
78 basin size and slope (Meade, 1996); d) increased sediment load causes increases in deposition in
79 the alluvial valley due to its buffering capacity (Walling, 2006); e) most of the increase in
80 sediment load is fine sediment (silt-clay) unless river flow increases (Owens et al., 2005); f) an
81 increased sediment load reaching the river mouth does not necessarily imply additional delta
82 growth since fine sediment is deposited in the ocean (prodelta) or in the delta plain (aggradation)
83 but does not contribute to delta fringe progradation (Orton and Reading, 1993); and g) dams

84 retain a significant portion of the river sediment load, though large-scale damming is a relatively
85 recent activity. Large dams store approximately 65 Gt/y, an Anthropocene magnitude that is
86 many times the total load transported by rivers (Syvitski et al., 2005).

87 Furthermore, reliable long-term records of sediment load dynamics are scarce (Whol et al.,
88 2015); and sediment record in deltas is often discontinuous; the limited C¹⁴ dates available
89 indicate that deposition is highly heterogeneous across space and time (Cearreta et al., 2016).
90 Consequently, it is difficult to test the contribution of land use changes on delta growth directly
91 (i.e., via field data linking additional delta growth to additional sediment erosion and transport).

92 Here we take an indirect approach to assessing the contribution of natural forcings and land-
93 use change to delta progradation, applying General Linear Models (GLMs) to a new database of
94 world deltas to quantify the contribution of these forcings. We hypothesized that no matter the
95 changes in the rhythm of delta growth, the final result (delta size) would be different if human
96 land-use changes have exacerbated sediment transport to the coast. We further hypothesized that
97 if human effects at the watershed scale were important to delta growth, a proportionally larger
98 delta size would be expected for river basins with a long history of human occupation and land
99 clearing. We measured delta size as the area included in the external contour of the emergent
100 delta as defined by the position of the delta fringe, so the conversion of marshes to open water in
101 the delta plain because of sea level rise and sediment deficit (Giosan et al., 2014) or the draining
102 of lagoons to create farmland are not considered in this analysis. We excluded from the analysis
103 the effects of direct human impacts in the delta plain that could influence delta size, such as
104 diking, artificial diversions of river distributaries, etc. Fourteen variables (out of 17) were
105 selected (Table 2) from 86 deltas and their basins worldwide (Fig. 1) to evaluate the effect of
106 watershed anthropogenic drivers versus climatic and physical drivers on delta size (the

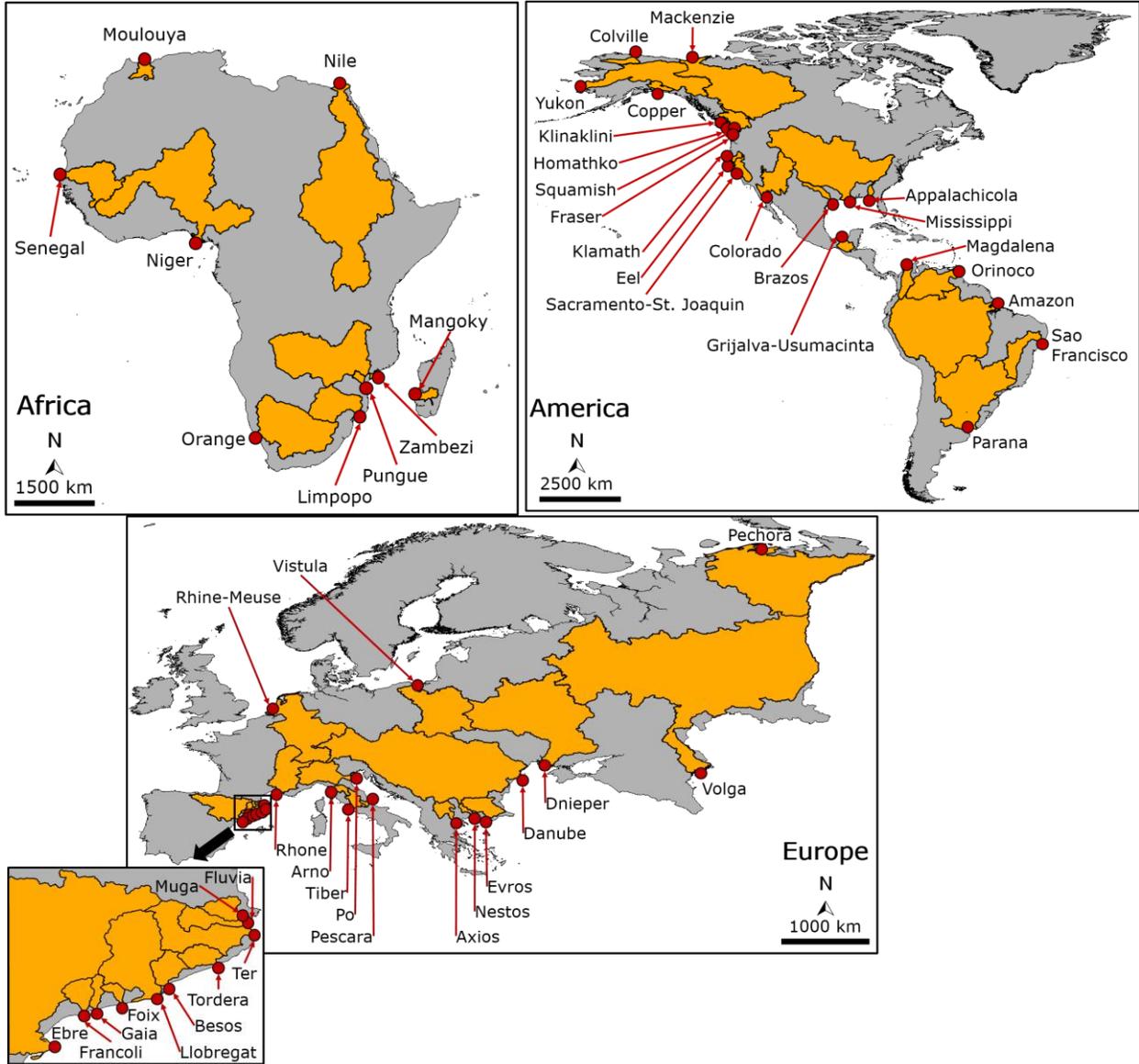
107 dependent variable) (See more details on methods and data sources on the corresponding
108 section).



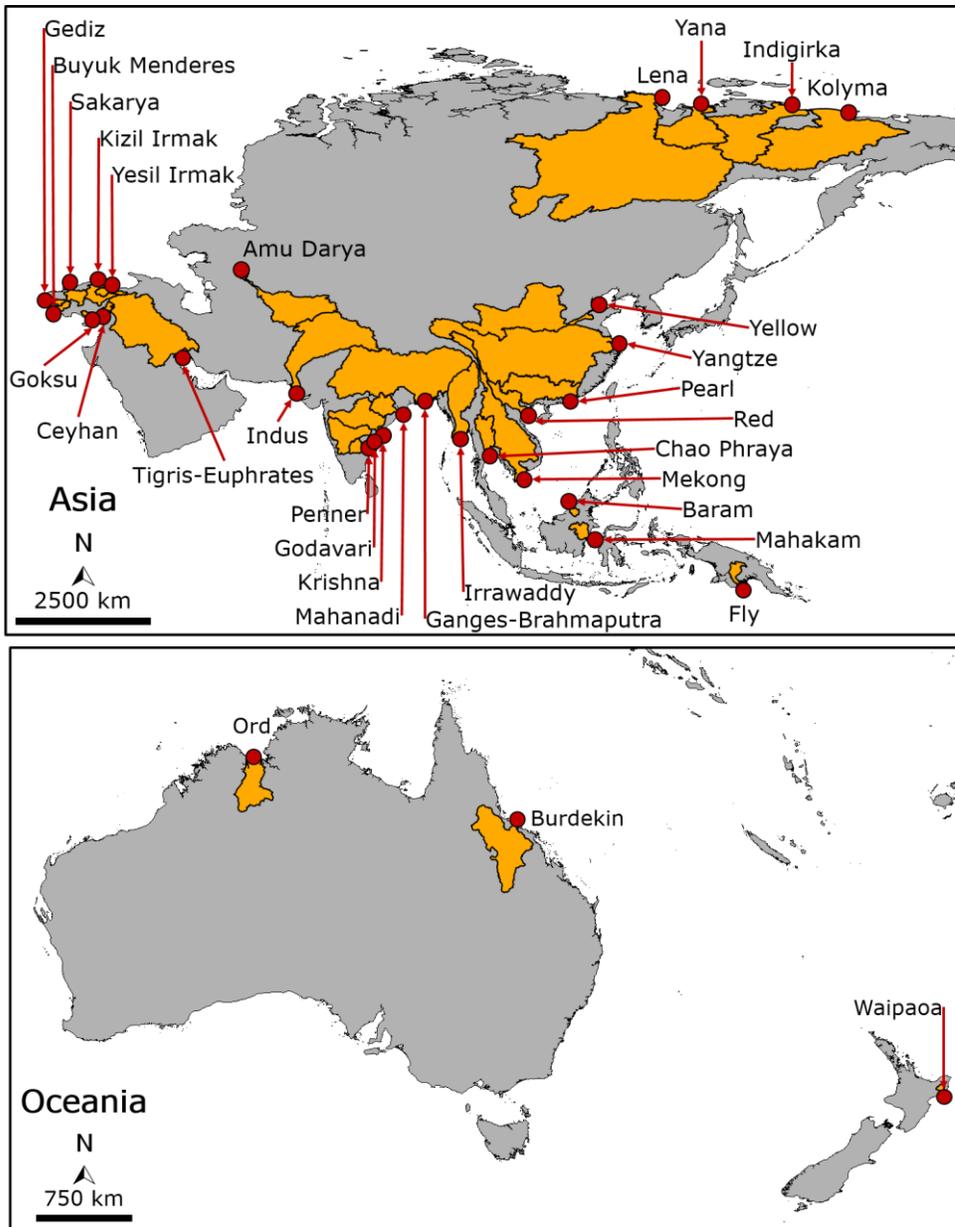
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110 **Figure 1.** World map showing the river basins covered in the present study.

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112 The new approach in this study is to include human-impact variables in the analysis together
113 with geomorphic and climatic forcings, using the new global dataset of deltas and their river
114 basins that is the most updated and complete to date. Another relevant contribution is the
115 separate analysis of Mediterranean and Black Sea deltas to determine if a stronger signature of
116 human impact on delta size is evident. We do not consider in the paper recent Anthropocene
117 effects that have been more investigated (i.e., impact of dams), rather we address the long-term
118 anthropic signature versus the geomorphic-climate signature on total delta size and sustainability.

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139 **Figure 2.** Maps showing the location and name of the studied deltas in the different continents.

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141 **Methods**

142 *Data collection*

143 Data was gathered for 86 deltas worldwide (29 of them opening into the Mediterranean and
144 Black Sea, named Med-Black deltas hereafter) by a systematic literature search of public
145 databases, scientific papers, and reports available from national and international literature. The
146 following groups of variables were initially considered for each delta: Location (latitude and
147 longitude); size and elevation (delta and basin surface area, river length and elevation at source,
148 river slope); mean river discharge, sediment load and sediment yield; mean rainfall; local marine
149 features (tidal amplitude and continental shelf slope); basin indicators of human impact:
150 population density, number of large cities, and number of dams; dam capacity; forest cover loss
151 and present land use (forest, grassland, savanna, and scrubland, wetlands, croplands and irrigated
152 croplands, urban-industrial, drylands). The variables, units and main sources of data collected are
153 summarized in Table 1.

154 We selected almost all world deltas down to ca. 100 km² (smaller for some Med-Black deltas),
155 excluding some cases for which available information was scarce. The final database used for the
156 multivariate analysis included 14 variables, among which were basin area and slope, flow and
157 sediment discharge, delta area, tidal range, shelf slope, land uses and vegetation cover, as well as
158 other human impacts (see Table 1). To keep the number of dependent variables as low as
159 possible to avoid excessive computing time (growing exponentially when one variable is added),
160 three variables were excluded from the initial 17. Delta area were obtained from published works
161 (e.g. Syvitski and Saito, 2007; Walker, 1998; Tockner et al., 2009; Kravtsova and Mit'kinykh,
162 2011) and the World Delta Database (Hart and Coleman, 2016). Measures from aerial

163 photographs and maps were used to validate and solve inconsistencies among databases. Mean
 164 river flow and sediment load were obtained from global databases such as Lehner et al. (2008),
 165 Fekete et al. (2002), and the Global River Sediment Yields Database (FAO, 2008); where
 166 possible, the early values less influenced by human activities were used. Most of the independent
 167 variables were extracted from public GIS databases with the Spatial Analyst Tool in ArcGIS 9.3
 168 (see Table 1 for source and resolution details), such as basin area, mean basin rainfall, mean
 169 population density in the year 2010, percent vegetation cover, and the human impact variables
 170 percent urban area, percent croplands, and percent forest deforestation. Mean basin slope was
 171 derived from the Global Multi-resolution Terrain Elevation Data 2010, and continental shelf
 172 slope (from the present river mouth to the outer limit of the shelf) from the General Bathymetric
 173 Chart of the Oceans (see Table 1). Tidal range was obtained from global long-term daily tidal
 174 charts, and dam capacity from the Global Reservoir and Dam Database; dam capacity was
 175 standardized by dividing capacity by basin area, thus obtaining relative dam capacity ($\text{Hm}^3 \text{ km}^{-2}$).
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177
 178 **Table 1.** Description of the variables for the analysis of 86 worldwide deltas. Data sources in
 179 numbers are cited in the main text. Delta area (dependent variable) was analyzed as function of
 180 14 independent variables (in bold) representing human and natural predictors.

Variable	Data acquisition	Source	Spatial Scale
Delta Area (km^2)	Bibliographic	Syvitski and Saito, 2007; Rakotomavo and Fromard, 2010; Walker, 1998; Tockner et al., 2009; Kravtsova and Mit'kinykh, 2011	

Basin Area (km²)	Derived from GIS	Peucker-Ehrenbrink, 2009; Meybeck and Ragu, 2012; Weatherall et al., 2015	15 ARCSEC 30 ARCSEC 30 ARCSEC
Mean Basin Slope (φ)	Derived from GIS	Revenga, 2003	7.5 ARCSEC
Mean River Flow (m³ s⁻¹)	Global databases	Hart and Coleman, 2016; Lehner et al., 2008; Fekete et al., 2002	
Sediment Load (kg s⁻¹)	Global databases	Hart and Coleman, 2016; Lehner et al., 2008; Fekete et al., 2002	
Tidal Range (m)	Global tidal charts		
Shelf Slope (φ)	Derived from GIS	Danielson and Gesch, 2011	30 ARCSEC
Mean Rainfall (mm)	GIS		30 ARCSEC
Dam Capacity Rate (Hm³ km⁻²)	Derived from GIS	Lehner, 2011	
Mean Population Density (# km⁻²)	GIS	CIESIN, 2016	30 ARCSEC
Forest Cover (%)	GIS	ESA, 2016	7.5 ARCSEC
Grassland, Savanna & Shrubland Cover (%)	GIS	ESA, 2016	7.5 ARCSEC
Urban Areas Cover (%)	Derived from GIS	DOD/USAF/AFWA	30 ARCSEC
Croplands Cover (%)	GIS	ESA, 2016	7.5 ARCSEC
Deforestation (%)	GIS	UNEP-WCMC, 1998	5 ARCMIN
Drylands Cover (%)	GIS	Cherlet et al., 2015	30 ARCSEC
Percent Habitat Lost (%)	GIS	Hoekstra et al., 2010	30 ARCSEC
Mean Human Footprint	GIS	Venter et al., 2016	30 ARCSEC

181

182 *Statistical analysis*

183 The effects of key environmental factors on delta size were analyzed using General Linear

184 Models (GLMs). An information-theoretical approach was applied to find the best approximating

185 models (Burnham and Anderson, 2002). GLMs were built including all possible combinations of

186 environmental features, excluding interactions, due to the large number of variables included.
187 The degree of support of each candidate model was assessed with the second order Akaike
188 Information Criterion (AICc), rescaled to obtain ΔAICc values ($\Delta\text{AICc} = \text{AICc}_i -$
189 minimumAICc). Candidate models were defined following two additional criteria: due to the
190 high correlation amongst the parameters, only models with a variance inflation factor (VIF) of \leq
191 3 were considered, to avoid multicollinearity effects in regression models (Maggini et al., 2006);
192 and model performance had to be significantly better than the null model (*i.e.* the model
193 including only the intercept), as judged by a likelihood-ratio test (Whittingham et al., 2005).
194 When the final set of candidate models is selected, models having ΔAICc values within 1–2 of
195 the best model have the most substantial support, those within 4–7 have considerably less
196 support, while models with $\Delta\text{AICc} > 10$ have essentially no support and might be omitted from
197 further consideration (Burnham and Anderson, 2002). No significant differences were observed
198 when $\Delta\text{AICc} = 2$ or $\Delta\text{AICc} = 7$ were considered, and the latter was used as supporting criterion.
199 The relative plausibility of each candidate model was assessed by calculating Akaike’s weights
200 (w_i), which range from 0 to 1, and can be interpreted as the probability that a given model is the
201 best model in the candidate set. Because no model was clearly the best one (*i.e.* $w_i \geq 0.9$), we
202 calculated model-average regression coefficients as the result of a weighted average (by model
203 w_i) of the regression coefficients across all models in which a given variable is present. The
204 relative importance of each independent variable was also calculated by the sum of w_i for all
205 models in which a given variable occurs (Burnham and Anderson, 2002). The more relative
206 importance the higher the values of selection probability (SP), which vary from 0 to 1. Model-
207 averaged regression coefficients (β) were compared with those from the full model to assess the
208 impact of model selection bias on parameter estimates (Maggini et al., 2006). Lower (in absolute

209 value) parameter bias indicates greater concordance between averaged and full models (reference
210 value 0.05; Genua-Olmedo et al. 2016). Higher positive or negative β coefficients (in absolute
211 value) indicate a stronger influence of an independent variable for differentiating the response
212 variable across all candidate models (Burnham and Anderson, 2002). For all of the candidate
213 models residuals were normally distributed according the Shapiro–Francia normality test ($P \geq$
214 0.25). All statistical analyses were performed with R software version 3.3.2; the MuMIn 1.15.6
215 package was used for multi-model inference analysis and Nortest 1.0-4 was used for normality
216 test analysis.

217

218 **Results**

219 Geomorphological and hydrological variables were most important in explaining delta size
220 (Table 2), especially continental shelf slope and mean river flow, were present in all the selected
221 models (SP= 1) and had the largest effect on the dependent variable ($\beta = -2.07$ and 0.84 ,
222 respectively), followed by sediment load (SP= 0.63, $\beta = 0.11$) and mean population density
223 (SP=0.56, $\beta =0.15$), both with a smaller effect. Mean rainfall and urban land cover had a small
224 negative effect, whereas the effects of deforestation, forest cover, grassland cover, cropland
225 cover, tidal range, dam capacity and basin slope were very small. However, the effect of
226 deforestation, forest cover and cropland area were more significant for the case of Med-Black
227 deltas, but still relatively small. Interestingly, mean rainfall showed a dual effect on delta size,
228 being positive for Med-Black deltas and negative for the other deltas. The weighted average
229 model selected 13 variables (out of 14), to predict delta size with a highly significant regression
230 coefficient (Fig. 2). Results from heavily human occupied basins of the Mediterranean and Black
231 Sea (Table 2, Figure S1) also showed a high regression coefficient between observed and

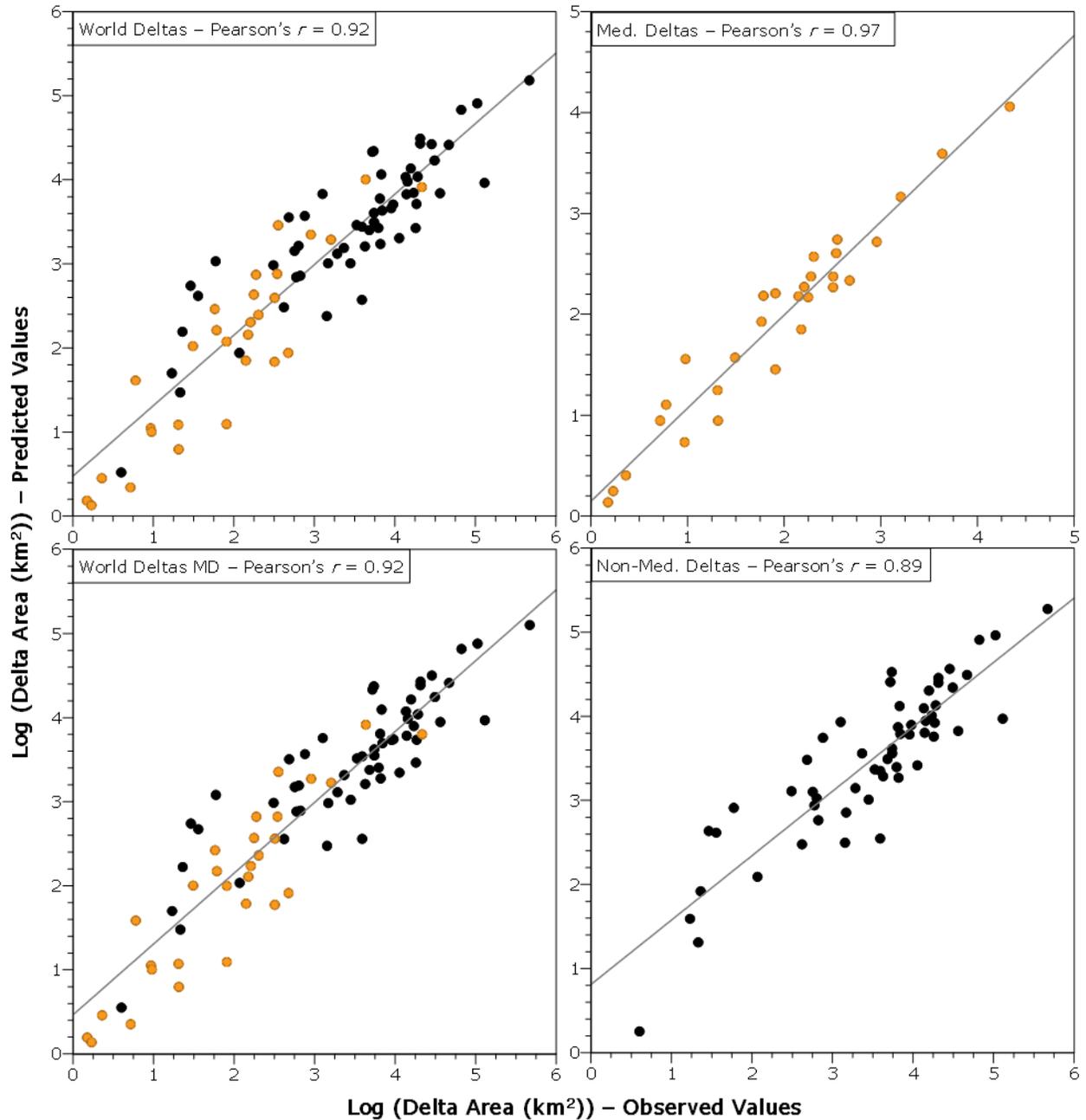
232 predicted values of delta size (Fig. 2). This model selected a total of 11 variables, excluding
233 sediment load, dam capacity and tidal range (Table 2). Interestingly, crop area and urban area
234 were selected in all models (SP=1) but they showed a moderate and negative effect on delta size
235 ($\beta = -0.14$ and -0.15 , respectively), whereas forest cover and deforestation were selected in fewer
236 models and showed a moderate positive effect ($\beta = 0.14$ and 0.13 , respectively). Physical
237 variables such as basin area and river slope were included in most models and had the largest
238 (positive) effect on delta size ($\beta = 0.83$ and 1.05 , respectively). In contrast to the model for world
239 deltas, Med-Black deltas show a very small influence of mean flow, shelf slope and population
240 density. Overall, human influence on the size of these deltas is more significant than in world
241 deltas, nevertheless physical and climatic drivers were the most influential. Noticeable, human
242 impacts show opposite or contradictory influences, since the negative effect of cropland area and
243 urban area on delta size counterbalance the positive effect of deforestation and forest cover
244 (which one might have expected to have an opposite influence). The net effect is that Med-Black
245 deltas do not appear to be a distinct group when plotted among the world deltas (Fig. 2), but they
246 tend to be smaller, as shown by further analysis using a new variable labeling Med-Black Deltas
247 (Table 2).

248

249 **Table 2.** Results from the information–theoretic framework analyses to evaluate the importance
250 of independent variables on delta size: World Deltas (86); Med. Deltas (29 Med-Black Deltas);
251 Non-Med. Deltas (57 World Deltas, excluding Med-Black Deltas); and World Deltas MD (86
252 World Deltas where Med-Black is included as factor). Model-averaged regression coefficients
253 (β), selection probability (SP), parameter bias, number (N) of candidate models and correlation
254 coefficient between observed and predicted values (r) are shown. Parameters included in the best

255 model are highlighted in grey. See *statistical analysis* for interpretation of model parameters (β ,
 256 SP and Bias).

Model Term	World Deltas <i>N</i> = 51, <i>r</i> = 0.92			Med. Deltas <i>N</i> = 8, <i>r</i> = 0.97			Non-Med. Deltas <i>N</i> = 144, <i>r</i> = 0.89			World Deltas MD <i>N</i> = 116, <i>r</i> = 0.92		
	SP	β	Bias	SP	β	Bias	SP	β	Bias	SP	β	Bias
Intercept	1.000	1.044	0.476	1.000	-5.995	-0.412	1.000	0.868	0.240	1.000	1.159	0.428
Mean River Flow (m ³ s ⁻¹)	1.000	0.837	0.212	0.145	0.121	0.329	1.000	0.910	0.099	1.000	0.794	0.266
Shelf Slope (ϕ)	1.000	-2.072	-0.041	0.116	-0.072	-2.299	1.000	-2.470	-0.128	1.000	-2.028	-0.031
Sediment Load (kg s ⁻¹)	0.634	0.111	0.040	<i>Not selected</i>			0.409	0.069	-0.636	0.701	0.135	0.014
Population (# km ⁻²)	0.559	0.152	-0.605	0.079	0.042	-13.607	0.108	0.010	-1.767	0.569	0.152	-0.656
Rainfall (mm)	0.556	-0.267	0.333	0.394	1.068	0.358	0.422	-0.215	0.511	0.534	-0.261	0.412
Urban Areas (%)	0.544	-0.044	-0.709	1.000	-0.149	-0.869	0.056	0.000	-75.990	0.481	-0.037	-0.723
Deforestation (%)	0.379	0.018	-1.574	0.527	0.131	-0.863	0.588	0.042	-0.865	0.390	0.019	-1.564
Forest Cover (%)	0.112	-0.003	-6.899	0.664	0.139	-1.388	0.430	-0.034	-1.624	0.088	-0.002	-8.111
GSS (%)	0.105	-0.003	-21.911	0.058	-0.003	27.178	0.543	-0.040	-1.996	0.126	-0.003	-15.137
Croplands (%)	0.090	0.002	30.082	1.000	-0.140	0.119	0.259	0.013	4.871	0.123	0.004	10.797
Tidal Range (m)	0.084	-0.014	-12.140	<i>Not selected</i>			0.143	-0.047	-6.025	0.343	-0.169	-2.118
Dam Capacity (Hm ³ km ⁻²)	0.070	0.001	-1.152	<i>Not selected</i>			0.055	0.000	-68.924	0.072	0.001	-3.572
Basin Slope (ϕ)	0.063	-0.001	123.641	0.767	1.046	-0.227	0.226	0.091	-3.414	0.054	0.002	-154.444
Basin Area (km ²)	<i>Not selected</i>			0.855	0.827	-0.133	0.112	0.030	-3.897	<i>Not selected</i>		
Med-Black Delta										0.558	-0.191	-1.232



257

258 **Figure 2.** Regression plots of World Deltas, Med-Black Deltas, World Deltas including Med-
 259 Black Deltas as a variable, and Non-Mediterranean Deltas, showing the Pearson's Correlation
 260 coefficients (r) of predicted versus observed values of the dependent variable (delta size)
 261 obtained from the General Linear Models (GLMs). Med-Black Deltas are shown in yellow color.

262

263 **Discussion**

264 Results clearly show that the main drivers of delta size at the watershed level are strongly
265 related to the hydrology (climate), and the morphology of the river basin and the continental
266 shelf. A previous study by Syvitski and Saito (2007) also found river discharge to be the main
267 explanatory variable of delta area, by using a scaling model including average discharge of its
268 feeder river, the total sediment load, and the shelf depth ($R^2 = 0.91$, 0% bias). In contrast, total
269 sediment load has a small effect in the model presented here, likely because land-use variables
270 are explicitly included in the analysis. Population density was found to be the most important
271 among the anthropogenic variables for the case of world deltas; (see also Syvitski and Milliman,
272 2007). Population density integrates several aspects of human interference in river basin
273 hydrology and sedimentary processes. Population density is often related to deforestation rates
274 (Maselli and Trincardi, 2013) and to a larger area devoted to farmland and urban areas. Perhaps
275 the reason population density is more important than other anthropogenic variables is that it
276 captures the interference of human activities on the alluvial valley and river hydro-
277 sedimentology. Aside from the more recent (20th century) effect of dams, which is not having a
278 significant effect on total delta size yet, the long-term alteration of the river course and its
279 floodplain such as narrowing and dredging, straightening of meanders, construction of levees,
280 occupation of the flood plain and the delta plain, among others, are reported to have a significant
281 impact on sediment transport and the buffering capacity of the flood plain; such as the case of
282 river basins undergoing millennial human occupation (Hori et al., 2001).

283 In general, results concerning world deltas and deltas other than Med-Black ones are more
284 consistent and clearer than those obtained for the Med-Black deltas. The main reason is likely
285 related to the small number of study cases (29) in relation to the analyzed variables (14), leading

286 also to a stronger influence of outliers; so we think the results of this subset must be interpreted
287 with more caution than those from the whole data set.

288 Contrary to world deltas, mean river flow and shelf slope are not important in explaining
289 Med-Black delta size. Here the variables having more influence on delta size are rainfall, basin
290 slope and basin area (again natural variables), whereas cropland area and urban area are the
291 human variables with more influence but with much lower model-averaged regression
292 coefficients (i.e. less influence on delta size). One possible explanation is that Mediterranean
293 river flow regime is flashier and has been strongly modified from long ago, so in this case mean
294 river flow is not so relevant to explain the capacity of sediment transport to the coast (and the
295 consequent delta growth). This may explain why basin slope and area, together with rainfall have
296 a larger influence in Med-Black deltas, and why sediment load is not selected by the model. Also
297 in contrast with world deltas, shelf slope is not important in determining Med-Black delta size,
298 possibly because the slope is more homogeneous among deltas. The negative effect of mean
299 rainfall on delta size for the case of world deltas is more difficult to explain; one possible
300 interpretation is that rainier river basins tend to have more regular flow regimes (proportionally
301 less flashy).

302 The small effect of deforestation and other land use changes on delta size is likely related to
303 erosion and sediment delivery to the coast being buffered in space and time within the basin, the
304 so called “sediment delivery problem” (Walling, 1999). The relatively more important role of
305 deforestation in Med-Black deltas could be explained because land clearing took place earlier
306 and the river basins are smaller (and with higher slope) than the world average. Many larger river
307 basins and deltas reflect long-term Holocene forces, where it is more difficult and takes longer
308 for hinterland perturbations to be manifested as coastal deposits (Syvitski and Milliman, 2007).

309 The large discrepancy between cropland soil losses estimated from field studies and modelling
310 (~75 Gt/yr) and sediment fluxes in rivers (~21 Gt/yr) is mostly explained by the deposition of
311 most eroded material immediately adjacent to its source; only 2% of the terrestrial surface of the
312 Earth is apparently enough to accumulate the sediment eroded from the remaining 35% devoted
313 to farmland (Wilkinson and McElroy, 2007). For example, a study in SE Minnesota (USA)
314 estimated that 47 to 65% of all sediment eroded in the past 137 years has moved no farther than
315 3 or 4 km (Beach, 1994). Another study in an area of the Magdalena river basin (Colombian
316 Andes) undergoing severe recent deforestation (241% increase) resulted in a 33% increase in
317 erosion rates that has translated into an increase of only 9% in total river sediment load (Restrepo
318 et al., 2015).

319 Moreover, subaerial delta progradation is mostly dependent on sand deposition (Nittrouer
320 and Viparelli, 2014) and subaqueous delta progradation is favored with a higher proportion of
321 finer sediment (Swenson et al. 2005), so the increase in sediment load due to human activity
322 mostly leads to an increase in fine sediment, which may be transported offshore in a large
323 proportion (Giosan et al., 2014; Orton and Reading, 1993) or deposited in delta plain wetlands
324 (Day et al., 2007). However, progradation rates of the subaerial and subaqueous deltas may differ
325 considerably besides grain size effects, since they are also influenced by the frequency and
326 magnitude of coastal storms, as well as by river discharge and flood frequency (Swenson et al.
327 2005). Therefore, deltas that prograde into low-energy marine environments fed by high-energy
328 rivers should be those in which delta size would be more influenced by an increase of fine
329 sediment inputs due to human alteration of the watershed.

330 Study results and existing literature do not suggest that the well documented anthropic
331 increase in sediment delivery to the coast has caused in most of cases a significant additional

332 delta progradation (higher than natural rates). In the Waipaoa River (New Zealand) sediment
333 load increased more than 6 times due to deforestation (Kettner et al., 2007) primarily being mud,
334 whereas delta progradation was largely controlled by sand; net accumulation of mud occurs
335 behind the prograding sandy shoreface (Wolinsky et al., 2010). In the Irrawaddy (Myanmar),
336 since 1850 less than 9% of the sediment load delivered to the delta has contributed to the
337 observed progradation (Hedley et al., 2010), while in the Mekong Delta (Vietnam) the rate of
338 delta front migration and stable sediment discharge during the last 3000 years indicate that a
339 large increase in sediment discharge due to human activities did not occur in this typical
340 Southeast Asian river (Ta et al., 2002). The subaerial delta of the Ganges-Brahmaputra is
341 prograding very slowly relative to its subaqueous delta (Allison, 1998), and regions of the
342 Amazon delta shoreface are eroding concomitant with rapid progradation of its subaqueous delta
343 (Nittrouer et al., 1996). A recent paper showed that the reduction in sediment supply to the
344 Mekong Delta is likely due to a shift in the tropical cyclone activity (Darby et al., 2016).

345 The recent increase in dam building is certainly affecting delta growth, but significant
346 changes in delta size (in relation to total surface) may take longer to manifest. Although dam
347 construction has dramatically reduced sediment fluxes to the coast (Syvitski et al., 2005), our
348 results show no effect of dam capacity on delta size possibly because it is a recent effect that is
349 not significant yet in global terms. However, in some river basins (especially in developing
350 countries) the near synchronicity of dam building with deforestation and other land use changes
351 limits the impact of upstream sediment load increases and could in part explain the small effect
352 of land use changes on delta size.

353 Recent studies in the Ebro Delta, considered a typical example of “man-made” delta
354 (Maselli and Trincardi, 2013), suggest that there has not been a large increase in sediment

355 delivery and delta growth due to human activities. The long-term increase in sediment load of the
356 lower Ebro River due to deforestation and human occupation of the basin has now been
357 estimated at around 40% (Xing et al., 2014), rather than the 8-fold increase suggested previously
358 (Walling, 2006). The former study (Xing et al., 2014) also showed that the background sediment
359 load of the Ebro River under low impact conditions (4000 years before present) was already high
360 (30 million tones yr⁻¹). Another study (Cearreta et al., 2016) has dated the origin of the present
361 Ebro Delta back to the Early Holocene, not to the Late Holocene (Maselli and Trincardi, 2013;
362 Guillén and Palanques, 1997) contradicting the hypothesis of the Ebro as “man-made” due to
363 watershed deforestation.

364 Results suggest that similar conclusions may hold for other Mediterranean and Asian deltas,
365 which have a long history of human influence but have not been created as consequence of river
366 basin land use changes. We propose that the concept of “human-made delta” should be used only
367 for the case of new deltas totally (or almost totally) created by human alteration of the river basin
368 greatly increasing sediment load. One of the few known candidates is the Mangoky Delta in
369 Madagascar, where the river mouth quickly prograded in the last 30 yr due to heavy
370 deforestation in the hinterland (Rakotomavo and Fromard, 2010). These cases essentially may
371 occur in the context of tropical river basins with heavy rains, high relief and highly erodible
372 soils. We consider human-made deltas as extreme examples of “human-influenced deltas”,
373 which are those in which their size, shape and other geomorphological features have been altered
374 by changes in the sediment flux due to land use changes in the watershed. This is the case for
375 most of the world’s deltas, except pristine deltas which are scarce nowadays. Then there is the
376 case of “human-engineered deltas”, those modified by the action of humans within the delta
377 itself, mostly due to diversion or diking of the river distributaries and the delta plain; good

378 examples of heavily engineered deltas are those of the Po and Yellow Rivers. In those cases
379 deltas existed under pristine conditions, but they were smaller, with a different shape, in a
380 different place, etc. Indeed, human-made large river diversions can create "man-made delta
381 lobes" (as is the case of the new Yellow Delta), since the human deviation of the river far away
382 of the previous mouth creates a new delta lobe that is not overlapping with the previous one.

383

384 **Conclusions**

385 This paper sheds light on the debate about the Anthropocene and the idea that some deltas
386 have formed or are much larger because of the human alteration of the river basin that leads to an
387 increase in sediment delivery to the coast. Some authors have even coined the concept of “man-
388 made deltas” and argued that Mediterranean deltas were mainly formed due to human alteration
389 of the basin. Our results contradict the idea of human-made deltas since delta size is essentially
390 determined by natural forcings, even in Mediterranean deltas.

391 One of the findings of special interest is the weak signature of deforestation on delta size.
392 Obviously human activity has increased the flux of sediments through land use changes (or in
393 some cases, to reduce it), but we argue that it is mostly fine sediments that contribute little to
394 delta progradation unless it is accompanied by a change in river flow that increases the capacity
395 of the river to transport sand. This is not to underestimate the adverse effects of human activities
396 on deltas, by development and reclamation, and by degradation of delta plain wetlands through
397 prevention of sediments reaching the delta plain due to levee and dam construction. A more
398 precise terminology is proposed to clarify concepts such as “human-made”, “human-engineered”
399 or “human-influenced” deltas.

400

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409

410 **Data availability**

411 All the original data are of public access and sources can be found in Table 1 (see references).
412 The analyzed database is available as supporting information.

413

414 **References**

- 415 Allison, M.A., Kuehl, S.A., Martin, T.C., Hassan A., 1998. Importance of flood-plain
416 sedimentation for river sediment budgets and terrigenous input to the oceans: Insights from
417 the Brahmaputra-Jamuna River. *Geology* 26 (2), 175–178.
- 418 Allison, M.A., 1998. Historical changes in the Ganges-Brahmaputra delta front. *J. Coastal Res.*
419 14, 1269–1275.
- 420 Beach, T., 1994. The fate of eroded soil: sediment sinks and sediment budgets of agrarian
421 landscapes in southern Minnesota, 1851–1988. *Ann. Assoc. Am. Geogr.* 84 (1), 5–28.
- 422 Burnham, K.P., Anderson, D.R., 2002. Model selection and multimodel inference: a practical
423 information–theoretic approach. Springer-Verlag, New York.

424 Center for International Earth Science Information Network (CIESIN)., 2016. Columbia
425 University. Gridded Population of the World, Version 4 (GPWv4): Population Density.
426 Palisades, NY: NASA Socioeconomic Data and Applications Center (SEDAC).
427 <http://dx.doi.org/10.7927/H4NP22DQ>

428 Cearreta, A., Benito, X., Ibañez, C., Trobajo, R., Giosan, L., 2016. Holocene
429 palaeoenvironmental evolution of the Ebro Delta (Western Mediterranean Sea): Evidence
430 for an early construction based on the benthic foraminiferal record. *Holocene* 26 (9), 1438–
431 1456.

432 Cherlet, M., Reynolds, J., Hutchinson, C., Hill, J., von Maltitz, G., Sommer, S., Fensholt, R.,
433 Horion, S., Shepherd, G., Weynants, M., Kutnjak, H., Smid, M., 2015. World atlas of
434 desertification, mapping land degradation and sustainable land management Opportunities.
435 The Joint Research Centre (JRC) of the European Commission in partnership with the
436 United Nations Environment Program (UNEP).

437 Danielson, J.J. Gesch, D.B., 2011. Global multi-resolution terrain elevation data 2010
438 (GMTED2010). U.S. Geological Survey Open-File Report 2011–1073 (26 pp.).

439 Darby, S.E., Hackney, C.R., Leyland, J., Kumm, M., Lauri, H., Parsons, D.R., Best, J.L.,
440 Nicholas, A.P., Aalto, R., 2016. Fluvial sediment supply to a mega-delta reduced by shifting
441 tropical-cyclone activity. *Nature* 539, 276–279.

442 Day, J.W., Boesch, D.F., Clairain, E.J., Kemp, G.P., Laska, S.B., Mitsch, W.J., Orth, K.,
443 Mashriqui, H., Reed, D.J., Shabman, L., Simenstad, C.A., Streever, B.J., Twilley R.R.,
444 Watson, C.C., Wells, J.T., Whigham, D.F., 2007. Restoration of the Mississippi Delta:
445 lessons from hurricanes Katrina and Rita. *Science* 315, 1679–1684.

446 DOD/USAF/AFWA. Air Force Weather Agency, U.S. Air Force, U.S. Department of Defense.
447 Nighttime Lights Annual Composites V4. <https://ngdc.noaa.gov/eog/download.html>
448 ESA (European Space Agency)., 2016. CCI Land Cover Product User Guide version 2.5, ESA
449 CCI LC project. <http://maps.elie.ucl.ac.be/CCI/viewer/index.php>
450 Fekete, B., Vörösmarty, C. Grabs, W., 2002. Global composite runoff fields on observed river
451 discharge and simulated water balances/ Water System Analysis Group, University of New
452 Hampshire, and Global Runoff Data Centre. Federal Institute of Hydrology, Koblenz,
453 Germany.
454 FAO (Food and Agriculture Organization of the United Nations)., 2008. AQUASTAT: Global
455 River Sediment Yields Database. Land and Water Development Division.
456 <http://www.fao.org/nr/Water/aquastat/sediment/index.stm>
457 Giosan, L., Syvitski, J., Constantinescu, S., Day, J., 2014. Protect the world's deltas. *Nature* 516,
458 31–33.
459 Giosan, L., Coolen, M.J., Kaplan, J.O., Constantinescu, S., Filip, F., Filipova-Marinova, M.,
460 Kettner, A.J., Thom, N., 2012. Early anthropogenic transformation of the Danube-Black Sea
461 system. *Sci. Rep.* 2 (582), 1–6.
462 Guillén, J., Palanques, A., 1997. A historical perspective of the morphological evolution in the
463 lower Ebro River. *Environ. Geol.* 30 (3), 174–180.
464 Hamilton, L.S., 1987. What are the impacts of Himalayan deforestation on the Ganges-
465 Brahmaputra lowlands and delta? Assumptions and facts. *Mt. Res. Dev.* 7 (3), 256–263.
466 Hart, G.F. Coleman, J., 2016. The World Deltas Database Framework. Louisiana State
467 University. <https://www.geol.lsu.edu/WDD>.

468 Hedley, P.J., Bird, M. I., Robinson, R.A.J., 2010. Evolution of the Irrawaddy delta region since
469 1850. *Geogr. J.* 176 (2), 138–149.

470 Hoekstra, J.M., Molnar, J.L., Jennings, M., Revenga, C., Spalding, M.D., Boucher, T.M.,
471 Robertson, J.C., Heibel, T.J., Ellison, K., 2010. *The Atlas of Global Conservation: Changes,*
472 *Challenges, and Opportunities to Make a Difference.* University of California Press,
473 Berkeley.

474 Hori, K., Saito, Y., Zhao, Q., Wang, P., Li, C., 2001. Progradation of the Changjiang River delta
475 since the mid-Holocene. *Sci. China Ser. B.* 44(1), 87–91.

476 Kettner, A.J., Gómez, B. Syvitski, J.P.M., 2007. Modeling suspended sediment discharge from
477 the Waipaoa River system, New Zealand: the last 3000 years. *Water Resour. Res.* 43 (7),
478 W07411.

479 Kravtsova, V.I. Mit'kinykh, N.S., 2011. Mouths of world rivers in the atlas of space images.
480 *Water Resour.* 38 (1), 1–17.

481 Lehner, B., Verdin, K., Jarvis, A., 2008. New global hydrography derived from space borne
482 elevation data. *EOS Trans. Am. Geophys. Union* 89 (10), 93-94.

483 Lehner, B., 2011. High resolution mapping of the world's reservoirs and dams for sustainable
484 river flow management. *Front. Ecol. Environ.* 9 (9), 494-502.

485 Maggini R., Lehmann A., Zimmermann N.E. Guisan A., 2006. Improving generalized regression
486 analysis for the spatial prediction of forest communities. *J. Biogeogr.* 33, 1729–1749.

487 Maselli, V. Trincardi, F., 2013. Man made deltas. *Sci. Rep.* 3, 1926.
488 <https://doi.org/10.1038/srep01926>.

489 Meade, R.H., 1996. River-sediment inputs to major deltas. In: Sea-level rise and coastal
490 subsidence. Springer, Dordrecht, pp. 63–85.

491 Meybeck, M. Ragu, A., 2012. GEMS-GLORI World River Discharge Database. Université Pierre
492 et Marie Curie, Paris, France. doi:10.1594/PANGAEA.804574.

493 Nittrouer, J.A. Viparelli, E., 2014. Sand as a stable and sustainable resource for nourishing the
494 Mississippi River delta. *Nat. Geosci.* 7 (5), 350.

495 Nittrouer, C.A., DeMaster, D. J., 1996. The Amazon shelf setting: Tropical, energetic, and
496 influenced by a large river. *Cont. Shelf Res.* 16, 553–573.

497 Notebaert, B., Verstraeten, G., Govers, G., Poesen, J., 2010. Quantification of alluvial sediment
498 storage in contrasting environments: Methodology and error estimation. *Catena* 82 (3), 169–
499 182.

500 Owens, P.N., Batalla, R.J., Collins, A.J., Gomez, B., Hicks, D.M., Horowitz, A.J., Kondolf,
501 G.M., Marden, M., Page, M.J., Peacock, D.H., Petticrew, E.L., Salomons, W., Trustrum,
502 N.A., 2005. Fine- grained sediment in river systems: environmental significance and
503 management issues. *River Res. Appl.* 21 (7), 693–717.

504 Orton G.J. Reading, H.G., 1993. Variability of deltaic processes in terms of sediment supply,
505 with particular emphasis on grain size. *Sedimentology* 40 (3), 475–512.

506 Peucker-Ehrenbrink, B., 2009. Land2Sea database of river drainage basin sizes, annual water
507 discharges, and suspended sediment fluxes. *Geoch. Geophys. Geosy.* 10 (6).

508 Rakotomavo, A., Fromard, F., 2010. Dynamics of mangrove forests in the Mangoky River delta,
509 Madagascar, under the influence of natural and human factors. *Forest Ecol. Manag.* 259 (6),
510 1161–1169.

511 Restrepo, J., Kettner, A.J. Syvitski, J.P.M., 2015. Recent deforestation causes rapid increase in
512 river sediment load in the Colombian Andes. *Anthropocene* 10, 13–28.

513 Revenga, C., 2003. Watersheds of the World CD-Rom. Washington, DC, World Resources
514 Institute.

515 Saito, Y., Yang, Z. Hori, K., 2001. The Huanghe (Yellow River) and Changjiang (Yangtze
516 River) deltas: a review on their characteristics, evolution and sediment discharge during the
517 Holocene. *Geomorphology* 41 (2), 219–231.

518 Shi, C., Zhang, L., Xu, J., Guo, L., 2010. Sediment load and storage in the lower Yellow River
519 during the late Holocene. *Geogr. Ann. A* 92 (3), 297–309.

520 Swenson, J. B., Paola, C., Pratson, L., Voller, V. R., Murray, B., 2005. Fluvial and marine
521 controls on combined subaerial and subaqueous delta progradation: Morphodynamic
522 modelling of compound-clinoform development. *J. Geophys. Res.* 110, F02013,
523 doi:10.1029/2004JF000265.

524 Syvitski, J.P., Vörösmarty, C.J., Kettner, A.J., Green, P., 2005. Impact of humans on the flux of
525 terrestrial sediment to the global coastal ocean. *Science* 308, 376–380.

526 Syvitski, J.P.M. Kettner A., 2011. Sediment flux and the Anthropocene. *Philos. T. Roy. Soc. A*
527 369 (1938), 957–975.

528 Syvitski, J.P.M. Saito, Y. Morphodynamics of deltas under the influence of humans. *Global*
529 *Planet. Change* 57 (3), 261–282.

530 Syvitski, J.P.M., Milliman, J.D., 2007. Geology, geography and humans battle for dominance
531 over the delivery of sediment to the coastal ocean. *J. Geol.* 115, 1–19.

532 Ta, T.K.O., Nguyen, V.L., Tateishi, M., Kobayashi, I., Tanabe, S., Saito, Y., 2002. Holocene
533 delta evolution and sediment discharge of the Mekong River, southern Vietnam. *Quaternary*
534 *Sci. Rev.*21 (16), 1807–1819.

535 Tockner, K., Uehlinger, U. Robinson, C.T., 2009. *Rivers of Europe*. Academic Press.

536 Walker, H.J., 1998. Arctic Deltas. *J. Coastal Res.* 14(3), 718–738.

537 UNEP-WCMC, Global Generalized 'Original' Forest dataset (V 1.0)., 1998. UNEP World
538 Conservation Monitoring Centre, Cambridge, UK Publisher. [https://www.unep-](https://www.unep-wcmc.org/resources-and-data/generalised-original-and-current-forest)
539 [wcmc.org/resources-and-data/generalised-original-and-current-forest](https://www.unep-wcmc.org/resources-and-data/generalised-original-and-current-forest)

540 Venter, O., Sanderson, E.W., Magrath, A., Allan, J.R., Beher, J., Jones, K.R., Possingham, H.P,
541 Laurance, W.F., Wood, P., Fekete, B.M., Levy, M.A., Watson, J.E.M., 2016. Global
542 terrestrial human footprint maps for 1993 and 2009. *Sci. Data* 3, 160067. doi:
543 10.1038/sdata.2016.67.

544 Walling, D.E., 2006. Human impact on land-ocean sediment transfer by the world's rivers.
545 *Geomorphology* 79, 192–216.

546 Walling, D.E., 1999. Linking land use, erosion and sediment yields in river basins. In: *Man and*
547 *River Systems*. Springer, Netherlands, pp. 223-240.

548 Waters, C.N., Zalasiewicz, J., Summerhayes, C., Barnosky, A.D., Poirier, C., Gałuszka, A.,
549 Hajdas, I., Cearreta, A., Edgeworth, M., Ellis, E., Ellis, M.A., Jeandel, C., Leinfelder, R.,
550 McNeill, J.R., Richter, D.B., Steffen, W., Syvitski, J., Vidas, D., Wagnreich, M., Williams,
551 M., Zhisheng, A, Grinevald, J., Odada, E., Oreskes, N., 2016, The Anthropocene is
552 functionally and stratigraphically distinct from the Holocene. *Science* 351 (6269), aad2622.

553 Whittingham M.J., Swetnam R.D., Wilson J.D., Chamberlain D.E. Freckleton R.P., 2005.
554 Habitat selection by yellowhammers *Emberiza citrinella* on lowland farmland at two spatial
555 scales: implications for conservation management. *J. Appl. Ecol.* 42, 270–80.

556 Wilkinson B.H. McElroy, B.J., 2007. The impact of humans on continental erosion and
557 sedimentation. *Geol. Soc. Am. Bull.* 119, 140–156.

558 Weatherall, P. Marks, K.M., Jakobsson, M., Schmitt, T., Tani, S., Arndt, J.E., Rovere, M.,
559 Chayes, D., Ferrini, V., Wigley, R., 2015. A new digital bathymetric model of the world's
560 oceans. *EarthSpace Sci.* 2(8), 331–345.
561 http://www.gebco.net/data_and_products/gridded_bathymetry_data/gebco_30_second_grid/

562 Wohl E., Bledsoe, B.P., Jacobson, R.B., Poff, N.L., Rathburn, S.L., Walters, D.M., Wilcox,
563 A.C., 2015. The natural sediment regime in rivers: broadening the foundation for ecosystem
564 management. *BioScience* 65 (4), 358–371.

565 Wolinsky, M.A., Swenson, J.B., Litchfield, N., McNinch, J.E., 2010. Coastal progradation and
566 sediment partitioning in the Holocene Waipaoa sedimentary system, New Zealand. *Mar.*
567 *Geol.* 270 (1), 94–107.

568 Xing F., Kettner, A.J., Ashton, A., Giosan, L., Ibáñez, C., Kaplan, J.O., 2004. Fluvial response to
569 climate variations and anthropogenic perturbations for the Ebro River, Spain in the last 4000
570 years. *Sci. Total Environ.* 473, 20–31.

571

572 **Author contributions**

573 C. Ibáñez is the lead author and contributed to writing the paper; C. Alcaraz built most of the
574 database, did the statistical analyses and contributed to the Material and Methods; N. Caiola

575 contributed to writing the paper; P. Prado contributed to the database and statistical analysis; R.
576 Trobajo contributed to the database and reviewed some parts of the paper; X. Benito contributed
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578 contributed to writing and reviewing the paper.